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6           **Uncertainty Quantification in Reconstruction of Sparse Water Quality Time Series:**  
7           **Implications for Watershed Health and Risk-Based TMDL Assessment**

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**Abstract**

21 Despite the plethora of methods available for uncertainty quantification, their use has been limited  
 22 in the practice of water quality (WQ) modeling. In this paper, a decision support tool (DST) that  
 23 yields a continuous time series of WQ loads from sparse data using streamflows as predictor  
 24 variables is presented. The DST estimates uncertainty due to residual errors using a relevance  
 25 vector machine. To highlight the importance of uncertainty quantification, two applications  
 26 enabled within the DST are discussed. The DST computes (i) probability distributions of four  
 27 measures of WQ risk analysis- reliability, resilience, vulnerability, and watershed health- as  
 28 opposed to single deterministic values and (ii) concentration/load reduction required in a WQ  
 29 constituent to meet total maximum daily load (TMDL) targets along with the associated risk of  
 30 failure. Accounting for uncertainty reveals that a deterministic analysis may mislead about the WQ  
 31 risk and the level of compliance attained with established TMDLs.

32 **Keywords: Decision support tool, Water quality risk analysis, TMDL, Relevance vector**  
 33 **machine, Uncertainty quantification, LOADEST**

34

**Software and data availability**

Name of the software	Web-based Decision Support Tool
Programming languages	MATLAB, JavaScript and PHP
Name of the Dataset available	<b>A Decision Support Tool for Water Quality Modeling</b>
Developer and contact information	Ganeshchandra Mallya, gmallya@purdue.edu
Year First available	2018
Software required	Any web browser
Availability	Through the URL <a href="https://engineering.purdue.edu/WaterDST/StandAloneTool/">https://engineering.purdue.edu/WaterDST/StandAloneTool/</a>
User manual	Available online through the ‘User Manual’ tab of the Decision Support Tool
Cost	Free

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## 38 **1. Introduction**

39 Environmental decisions are often based upon mathematical models of the system under  
40 consideration (Beven, 2007; Refsgaard et al., 2006). For example, in total maximum daily load  
41 (TMDL, developed by United States Environmental Protection Agency, USEPA) development,  
42 models such as SWAT (Soil and water assessment tool, Arnold and Allen, 1999) or HSPF  
43 (Hydrological Simulation Program Fortran; Jia and Culver, 2006) are frequently used for  
44 simulation of water quality (WQ) constituents (e.g., Indiana Department of Environmental  
45 Management, IDEM, 2017). Typically, the parameters of a model are calibrated against available  
46 observations, and many observations are required to calibrate a complex model like SWAT. In  
47 some applications, continuous time series of streamflow observations and WQ constituent  
48 concentrations are required to assess the health of an impaired water body. Examples include (a)  
49 identification of sources of pollution in a waterbody (Mallya et al., 2018) and (b) computation of  
50 load reduction required (LRR) to restore a waterbody to healthy conditions (Park et al., 2015).  
51 Whereas streamflows are measured frequently in a watershed, WQ constituent such as suspended-  
52 solids, nitrogen, and phosphorus concentrations are measured sparsely (e.g., biweekly and only in  
53 summer months). Sparse WQ data are not amenable to direct use in reliable decision making  
54 (Kjeldsen and Rosbjerg, 2004). Therefore, various models have been developed for temporal  
55 *reconstruction* of WQ data (e.g., load estimator (LOADEST) Runkel et al., 2004). In this study,  
56 reconstruction refers to the estimation of WQ constituent concentration/load values at time-steps  
57 where WQ data is unavailable using the observed WQ constituent data.

58 Any mathematical representation of an open system, such as the ones encountered in  
59 environmental modeling, incur uncertainties due to lack of complete knowledge about the system,  
60 inadequate representation of dominant processes through mathematical equations, erroneous data  
61 used for parameter estimation, and difficult-to-represent local characteristics of the system (Beven,  
62 2007). These uncertainties in the modeling process should be considered to make informed  
63 management decisions (Beven, 2007). The quantification of uncertainties in hydrologic and WQ  
64 modeling is carried out using probabilistic methods (see Ahamadisharaf et al., 2019 for a review).  
65 In a typical probabilistic analysis, the residual time series (the difference between observed and  
66 simulated response of the system) is assumed to follow a probability distribution. The parameters  
67 of the probability distribution are estimated against observed residuals time series, and,  
68 subsequently, the calibrated probability distribution is used for uncertainty quantification. The

69 residual time series is an aggregate of *measurement errors*, *structural errors*, and errors in the  
70 numerical implementation of the model. Measurement errors refer to errors in the measurement of  
71 streamflow and WQ data. Structural errors exist because a model is an approximation of reality.  
72 Calibrated model parameters also incur uncertainty which is referred to as parametric uncertainty.  
73 Parametric uncertainty exists due to measurement errors, structural errors, and limited information  
74 in the data to calibrate the parameters. A modeler should ensure that errors in the numerical  
75 implementation are negligible. Thus, the residual time series is an aggregate of measurement and  
76 structural errors. Despite growing awareness about the importance of uncertainty due to structural  
77 errors (Brynisdottir and O’Hagan, 2014), measurement errors (Baldassarre and Montanari, 2009),  
78 unknown parameters (Melching and Bauwens, 2001) and residual errors (Beven and Binley, 1992;  
79 Borsuk and Stow, 2000; Borsuk et al., 2002; Chaudhary and Hantush, 2017; Hoque et al., 2012;  
80 Hantush and Chaudhary, 2014), uncertainty is rarely quantified in *practice* of WQ modeling. For  
81 example, in TMDL applications, the current practice is to use a margin of safety (MOS) to account  
82 for uncertainty in the relationship between the pollutant load and the quality of the receiving water  
83 body (Novotny, 2002). The MOS is typically assigned by making conservative assumptions or  
84 specified explicitly as a percentage (e.g., 5-10%) of the TMDL (NRC, 2001). Recently, Nunoo et  
85 al. (2020) found that, in 84% of the 37,841 TMDLs reported, uncertainty analysis was not carried  
86 out to select a margin of safety (MOS). Subjective or arbitrary specification of MOS might lead to  
87 overly conservative estimates and increased cost of implementation of pollution control measures  
88 (Zhang and Yu, 2004).

89 In WQ modeling, the pervasiveness of uncertainty has been long recognized (Beck, 1987) followed  
90 by several efforts to quantify it (e.g., Ahamadisharaf and Benham, 2020; Borsuk, 2003; Chaudhary  
91 and Hantush, 2017; Hoque et al., 2012; Jia and Culver, 2008; Reckhow, 2003; Shirmohammadi et  
92 al., 2006; Zhang and Yu, 2004; Zheng et al., 2011; Zheng and Han, 2016; and Zheng and Keller,  
93 2008). Uncertainty quantification for complex models tend to be complicated, time-consuming,  
94 and computationally demanding (see Smith et al., 2014, chap. 2 for some examples). Thus,  
95 researchers have sought simpler statistical models for simulation of WQ time series (e.g.,  
96 LOADEST).

97 For both physical and statistical models, a rich theory has been developed to quantify uncertainty  
98 by residual analysis (e.g., Smith et al., 2015). However, uncertainty quantification is frequently

99 avoided in practice for the following reasons (Pappenberger and Beven, 2006): (1) it is subjective,  
100 too difficult to perform and cannot be incorporated into decision making; (2) it is not required if  
101 one uses physically realistic models; and (3) it is too difficult for policy-makers to understand and  
102 does not really matter in making the final decision. Pappenberger and Beven (2006) further argued  
103 that the reasons cited above are untenable, and they emphasized the importance of an open  
104 discourse of uncertainty in environmental models. Reckhow (2003) pointed out that modelers  
105 should clearly communicate the uncertainties associated with their models to decision-makers. A  
106 solution to the problem ‘uncertainty analysis is too difficult to perform’ is availability of easy-to-  
107 use software packages that can be used by practitioners with little effort (e.g., Gronewold and  
108 Borsuk, 2009). In this paper, we present one such software through a decision support tool (DST).

109 The DST reconstructs WQ constituent time series by using streamflow values as predictor  
110 variables and employing the state-of-the-art relevance vector machine (RVM; Tipping, 2001) that  
111 can accommodate nonlinear transformations between streamflow and WQ data. Moreover, the  
112 RVM provides uncertainty estimates that are conditioned on streamflows and can account for  
113 errors in streamflows (not explored here). The DST uses the reconstructed WQ time series along  
114 with the uncertainty estimates for the following two applications:

115 (1) WQ risk assessment by computing indices such as reliability, resilience, vulnerability, and  
116 composite watershed health index (Hoque et al., 2012, Mallya et al., 2018).

117 (2) Computation of LRR of a WQ constituent so that the TMDL criterion is met with an  
118 acceptable risk of noncompliance (Camacho et al., 2018).

119 Specifically, the DST provides a probabilistic estimate of watershed health and the required  
120 reduction in pollutant concentration/load as a function of the risk of violating TMDL criteria. Fig.  
121 1 shows the overview of the three tasks carried out by the DST. In this study, using the St. Joseph  
122 River Watershed (spread over parts of Indiana, Michigan, and Ohio) as a test case, the two  
123 applications illustrate the importance of uncertainty in assessing watershed health and in  
124 conducting TMDL studies. While other investigators (e.g., Borsuk et al., 2002; Hantush and  
125 Chaudhary, 2014; and Camacho et al., 2018) have demonstrated the benefits of probabilistic  
126 uncertainty estimation in TMDLs, no study has implemented such a framework at the watershed  
127 scale, to the best of our knowledge.

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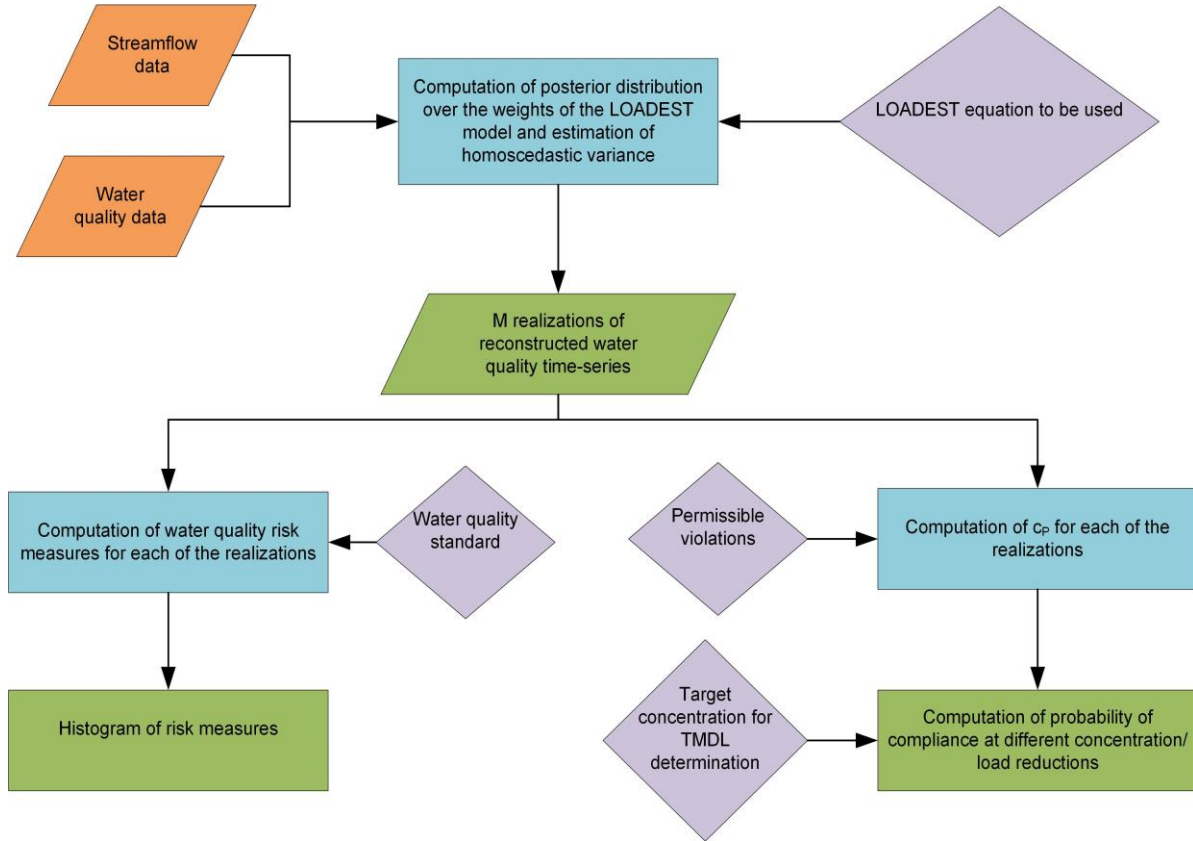


Fig. 1. A flowchart of the steps implemented in the Decision Support Tool

## 2. Theory

### 2.1 Reconstruction of water quality time series

Traditionally, simple regression equations LOADEST (Load Estimator, Runkel et al., 2004) have been used for reconstruction of WQ time series using available streamflows as predictors. However, the uncertainty associated with these estimates is rarely reported (or used). The DST uses RVM (Bishop, 2006 and Tipping, 2001) to estimate the uncertainty associated with the reconstruction. A general statistical model is used in the DST:

$$y_j = \sum_{i=1}^N w_i \phi_i(x_j) + \epsilon_j, \quad j = 1, 2, \dots, n \quad (1)$$

where  $y_j$  is log of load of WQ constituent at  $j^{th}$  time-step,  $\phi_i$  is the  $i^{th}$  (linear or non-linear) basis function,  $x_j$  is streamflow at  $j^{th}$  time-step,  $N$  is the number of the basis functions,  $w_i$  is the weight of the  $i^{th}$  basis function,  $\epsilon_j$  is the error in the estimation of  $y_j$  using  $\sum_i w_i \phi_i(x_j)$  as a model, and  $n$  is the number of time-steps at which observed streamflows are available. The basis functions

142 may be chosen to achieve a nonlinear transformations of streamflows into WQ constituents  
 143 (Tripathi and Govindaraju, 2007). The current version of the DST uses the same transformations  
 144 as the basis functions as are used in LOADEST due to the historical use of LOADEST in WQ  
 145 literature and to enable a comparison with LOADEST. For example, in case of the LOADEST  
 146 second equation (as listed in Table 1), the following three basis functions would be required: 1,  $\ln x$   
 147 and  $\ln x^2$ .

148 The  $\epsilon_j$  in Eq. (1) is assumed to be a zero-mean, Gaussian random variable with a homoscedastic  
 149 variance  $\sigma_\epsilon^2$ . Additionally, in RVM, each weight  $w_i$  is assigned a zero-mean Gaussian prior with  
 150 variance  $\alpha_i^{-1}$ . This specification of prior allows automatic determination of only the relevant basis  
 151 functions in Eq. (1) leading to predictions that are potentially robust to errors in predictor variables.  
 152 Subsequently, the posterior distribution of weight vectors  $\mathbf{w}$  (conditioned on the vector  $\mathbf{y}$ , matrix  
 153  $\Phi = [\phi(x_1), \phi(x_2), \dots, \phi(x_n)]$  of predictors, parameters  $\alpha_i$ 's, and the variance  $\sigma_\epsilon^2$ ) is computed  
 154 using the Bayes theorem. The posterior distribution of  $\mathbf{w}$  is found to be Gaussian with mean  $\boldsymbol{\mu}_w$   
 155 and covariance-matrix  $\Sigma_w$  such that

$$\begin{aligned}\Sigma_w &= (\sigma^{-2}\Phi\Phi^T + A)^{-1}, \\ \boldsymbol{\mu}_w &= \sigma^{-2}\Sigma_w\Phi\mathbf{y}, \text{ and} \\ A &= \text{diag}(\alpha_1, \alpha_2, \dots, \alpha_N).\end{aligned}\tag{2}$$

156 The predictive distribution at time-step  $t$ , given streamflow  $x_t$  is then Gaussian with mean

$$\mu_{y_t} = \boldsymbol{\mu}_w^T \phi(x_t),\tag{3}$$

157 and covariance between predicted log of load at  $t_1$  and  $t_2$

$$\Sigma_{t_1, t_2} = \sigma_\epsilon^2 \delta(t_1 - t_2) + \phi(x_{t_1})^T \Sigma_w \phi(x_{t_2}),\tag{4}$$

158 where  $\delta(t_1 - t_2)$  is the Dirac-delta function. Equation (4) shows that the uncertainty in  
 159 reconstruction at time-step  $t$  is dependent upon predictor  $x_t$  and covariance matrix  $\Sigma_w$  of weight  
 160 vector  $\mathbf{w}$ . For convenience in applications, the posterior distribution over the parameters  $\alpha_i$ 's and  
 161  $\sigma_\epsilon^2$  is approximated by the Dirac-delta function  $\delta(\boldsymbol{\alpha}_{MAP}, \sigma_{\epsilon, MAP}^2)$ , where  $\boldsymbol{\alpha}_{MAP}$  and  $\sigma_{\epsilon, MAP}^2$  are the  
 162 maximum posterior estimates of these parameters.

163 The DST requires daily streamflow time series and measurements of the WQ constituent as inputs  
 164 by the user. If the streamflow gauge and WQ monitoring stations are not co-located, it uses the  
 165 watershed area-ratio method (section 2.4) for the estimation of streamflow at the WQ monitoring



166 station. It allows the users to select any one of the nine LOADEST equations (Table 1), or to pick  
 167 the best LOADEST equation based on Akaike Information criterion (AIC; Akaike, 1973) if  
 168 desired. To represent the uncertainty in reconstructed WQ time series, it draws 10000 Monte Carlo  
 169 (MC) samples from the logarithm of load estimated as a multivariate Gaussian distribution (Fang  
 170 and Zhang, 1990). The mean and covariance matrix of the Gaussian distribution are given by Eqs.  
 171 (3) and (4), respectively. The 10000 MC simulations were found sufficient to obtain stable  
 172 estimates of lower and upper bounds of 90% and 95% prediction intervals/credible regions, WQ  
 173 risk measures (section 2.2), and TMDL compliance plots (section 2.3). Subsequently, the DST  
 174 computes 90% and 95% credible regions as follows. The 90% credible region is the region  
 175 bounded by 5<sup>th</sup> and 95<sup>th</sup> percentiles of MC samples; the percentiles are computed at each time-  
 176 step. The 95% credible region is the region bounded by 2.5<sup>th</sup> and 97.5<sup>th</sup> percentiles of MC  
 177 samples.

178 Even though DST uses the same basis-functions in RVM as those in LOADEST, a significant  
 179 difference between the RVM (as employed in the DST) and LOADEST (as employed by Park et  
 180 al., 2015) exists in the parameter estimation method. LOADEST uses adjusted-maximum-  
 181 likelihood estimation (AMLE) to estimate the weight vector  $\mathbf{w}$ . In RVM, the choice of the prior  
 182  $N(0, \alpha_i^{-1})$  on  $i^{th}$  weight  $w_i$  expresses a preference for smaller weights (Tripathi, 2009, pp. 13).  
 183 The smaller weights dampen observation errors in predictors as in LASSO and ridge regression  
 184 (Friedman et al., 2001). In some cases, the parameter  $\alpha_i^{-1}$  may converge to zero thus eliminating  
 185 the  $i^{th}$  basis function from the set of predictor variables; the reduced set of predictor variables  
 186 results in a computationally efficient model (Bishop, 2006 and Tipping, 2001).

187 Table 1. LOADEST equations

LOADEST equations
1. $\ln L = w_0 + w_1 \ln Q$
2. $\ln L = w_0 + w_1 \ln Q + w_2 \ln Q^2$
3. $\ln L = w_0 + w_1 \ln Q + w_2 \delta t$
4. $\ln L = w_0 + w_1 \ln Q + w_2 \sin(2\pi\delta t) + w_3 \cos(2\pi\delta t)$
5. $\ln L = w_0 + w_1 \ln Q + w_2 \ln Q^2 + w_3 \delta t$
6. $\ln L = w_0 + w_1 \ln Q + w_2 \ln Q^2 + w_3 \sin(2\pi\delta t) + w_4 \cos(2\pi\delta t)$
7. $\ln L = w_0 + w_1 \ln Q + w_2 \sin(2\pi\delta t) + w_3 \cos(2\pi\delta t) + w_4 \delta t$
8. $\ln L = w_0 + w_1 \ln Q + w_2 \ln Q^2 + w_3 \sin(2\pi\delta t) + w_4 \cos(2\pi\delta t) + w_5 \delta t$
9. $\ln L = w_0 + w_1 \ln Q + w_2 \ln Q^2 + w_3 \sin(2\pi\delta t) + w_4 \cos(2\pi\delta t) + w_5 \delta t + w_6 \delta t^2$

$\delta t$  = time in decimal units since first Julian of a year; for example, 31 Jan, 2005 is represented as 2005.085  
 $L$  = load (Kg day<sup>-1</sup>)  
 $Q$  = streamflow in m<sup>3</sup>s<sup>-1</sup> or ft<sup>3</sup>s<sup>-1</sup>

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## 188 2.2 Water quality risk analysis

189 The health of a watershed is quantified by using the following three measures (Hoque et al., 2012):  
190 reliability (R), resilience (R), and vulnerability (V). Additionally, a composite watershed health  
191 measure can be computed as a function of R-R-V (Mallya et al., 2018). Suppose,  $Y_t$  is the  
192 concentration or load of the reconstructed WQ constituent at time-step  $t$  with standard numerical  
193 target denoted by  $Y_t^*$  (concentration or load). The reliability ( $\rho$ ) is defined as the probability of the  
194 waterbody being in the compliant state, that is,

$$\rho = P\{Y_t \in S\} = 1 - P\{Y_t \in F\}, \quad (5)$$

195 where  $P\{\bullet\}$  denotes the probability of the event  $\{\bullet\}$ ,  $S$  denotes the event  $\{Y_t \leq Y_t^*\}$  denoting the  
196 safe/compliant state, and  $F$  denotes the event  $\{Y_t > Y_t^*\}$  denoting failed/noncompliant state. The  
197 definitions of the compliant and non-compliant state will be reversed in case of dissolved oxygen,  
198 that is,  $S = \{Y_t \geq Y_t^*\}$  and  $F = \{Y_t < Y_t^*\}$ . Given the time series of WQ constituent, the DST  
199 estimates  $\rho$  as

$$\rho = 1 - \frac{1}{n} \sum_{t=1}^n z_t, \quad (6)$$

200 where  $z_t = 1$ , when  $Y_t \in F$  and  $z_t = 0$ , when  $Y_t \in S$  and  $n$  is the total number of time-steps.  
201 Resilience ( $r$ ) is defined as the probability of the system to recover from a non-compliant state,  
202 that is,

$$r = P\{Y_{t+1} \in S | Y_t \in F\}, \quad (7)$$

203 and can be estimated as

$$r = \frac{\sum_{t=1}^n u_t}{\sum_{t=1}^n z_t}, \quad (8)$$

204 where  $u_t = 1$  when  $Y_{t+1} \in S$  and  $Y_t \in F$  and 0 otherwise, and  $z_t$  as defined above. In summary,  
205 a waterbody (or watershed) is resilient if it returns from a non-compliant state to a compliant state;  
206 the longer the waterbody takes to reach a compliant state from a non-compliant state, the less  
207 resilient the waterbody.

208 Vulnerability is defined as the magnitude of severity of violation during a noncompliant event.  
209 For WQ violations in a waterbody, there is no universal measure to quantify the severity of a

210 violation. Mallya et al. (2018) proposed a new measure referred to as robustness – opposite of  
 211 vulnerability – that scales between 0 and 1- as

$$v_o = \left\{ \prod_{t=1}^n \left( \frac{Y_t^*}{Y_t} \right)^{H[Y_t - Y^*]} \right\}^{\frac{1}{m}}, \quad (9)$$

212 where  $m$  is number of time-steps at which  $Y_t > Y_t^*$ ,  $H[\bullet]$  is the Heaviside function so that (9)  
 213 accounts only for the noncompliant events. When the deviations of  $Y_t$  from  $Y^*$  are large then  $v_o \rightarrow$   
 214 0; when deviations are small then  $v_o \rightarrow 1$ , which is consistent with definitions for reliability ( $\rho$ )  
 215 and resilience ( $r$ ). Vulnerability ( $v$ ) can now be defined as:

$$v = 1 - v_o \quad (10)$$

216 A composite measure of watershed health ( $h$ ) is defined as (Mallya et al., 2018):

$$h = (\rho r v_o)^{\frac{1}{3}} \quad (11)$$

217 Clearly, if  $\rho = r = v_o = 1$  then  $h = 1$ , i.e., the drainage area is healthy with respect to the WQ  
 218 constituent of interest. Similarly, if any one of the risk-measures is 0 then  $h = 0$ , i.e., the drainage  
 219 area is impaired with respect to WQ constituent of interest.

220 Reliability, resilience and vulnerability can be used to design appropriate measures to improve the  
 221 WQ of a waterbody. For instance, reliability should be used as a guiding measure if frequent  
 222 violations of a WQ constituent are not allowed. In cases where durations of violations are more  
 223 consequential than the frequency of violations, resilience is a useful measure. Similarly,  
 224 vulnerability is a useful measure when the goal is to reduce the severity of violations. Moreover,  
 225 spatial distribution of these WQ risk-measures over different streams can be used to identify the  
 226 critical sources of pollution in a watershed (e.g., Mallya et al., 2018). The usefulness of different  
 227 WQ risk measures also depend upon the timescale of analysis. For example, in some states, E. Coli  
 228 concentration should be below 235 cfu/100ml at least 89.5% of the time-steps at daily timescale  
 229 and should be below 126 cfu/100ml at 100% of time-steps at monthly timescale (Ahmadisharaf  
 230 and Benham, 2020). Thus, in case of E. Coli, watershed managers need to ensure that waterbody  
 231 has 0.895 and 1.00 reliability at daily and monthly timescales, respectively. Current version of the  
 232 DST computes WQ risk measures at the daily timescale. A future version will include WQ-risk  
 233 analysis at other timescales.

234 Uncertainty associated with reconstructed WQ data will carry over to the risk measures also. The  
235 DST uses MC method to estimate probability distributions of the R-R-V and the WH measures. It  
236 computes R-R-V and WH for each realization of reconstructed WQ time-series, and, subsequently,  
237 constructs the histograms of these measures from the ensemble values, and tabulates the mean and  
238 standard-deviation of the measures. The DST has a Google map interface which shows the  
239 locations of USGS WQ stations and National Water Quality Assessment (NAWQA) stations; a  
240 user can click on any one these WQ stations and fill out a form to reconstruct WQ data and compute  
241 WQ risk measures.

### 242 **2.3 Risk-based total maximum daily load (TMDL) analysis**

243 In this section, we present the theory behind the determination of the LRR to meet TMDL targets.  
244 Currently, the MOS is used to account for uncertainties associated with TMDL development. For  
245 example, suppose that the concentration of total phosphorus (TP) corresponding to a TMDL is  
246  $0.08 \text{ mg L}^{-1}$  and a deterministic model predicts that to maintain this concentration in a waterbody,  
247 the maximum allowable load in the watershed is  $600 \text{ Kg day}^{-1}$ . Then an arbitrarily selected 10%  
248 of the allowable load can be reserved for MOS. The 10% MOS serves to acknowledge that the  
249 magnitude of uncertainties in the model simulated TP are such that  $540 \text{ Kg day}^{-1}$  in watershed  
250 may also correspond to maximum TP concentration of  $0.08 \text{ mg L}^{-1}$ . However, the protective  
251 cushion provided by the 10% of the allowable load remains at best unknown. A small MOS may  
252 result in non-attainment of the water quality standard, but a large MOS can be inefficient and costly  
253 (Novotny, 2003). Therefore, a realistic estimate of uncertainty is required. The DST accommodates  
254 the uncertainty through MC simulations that yield an ensemble of  $M$  realizations of the WQ time  
255 series.

256 Borsuk et al. (2002) presented a framework for a probabilistic TMDL development. Let  $C^*$  be the  
257 user-defined target TMDL concentration of a waterbody, and suppose that the waterbody violates  
258 the TMDL standard at most  $p$  fraction of the times during the period of analysis. Then, one can  
259 define the probability of compliance ( $\kappa_p$ ) to permissible violations  $p$  as:

$$\kappa_p = P\{C_p \leq C^*\} = F_{C_p}(C^*), \quad (12)$$

260 where  $C_p$  denotes the concentration value that is exceeded  $p$  fraction of the times in a realization  
261 of WQ time series;  $p$  is the permissible fraction of violations (0.05, 0.10, ...etc.). The quantity  $C_p$

262 is a random variable that represents uncertainty associated with reconstructed WQ time series. The  
 263 quantity  $\kappa_p$  is the fraction of WQ constituent time series generated by MC method that are  
 264 compliant. The definition of  $\kappa_p$  would be reversed in case of dissolved oxygen, i.e.  $\kappa_p =$   
 265  $P\{C_p \geq C^*\}$ . The distribution of  $C_p$  is determined by MC method as follows.  
 266 For a reconstructed WQ constituent time series, a value of  $c_p$  is determined using

$$c_p = G^{-1}(1 - p), \quad (13)$$

267 where  $c_p$  is the  $100(1 - p)^{th}$  percentile of a reconstructed WQ time series,  $G$  is the empirical  
 268 cumulative distribution function (CDF) of reconstructed WQ constituent time series, and  $p$  is the  
 269 fraction of permissible violations (0.05, 0, 0.1, 0.15, 0.20 etc.). The quantity  $c_p$  may be interpreted  
 270 as threshold value to be compared against the target TMDL concentration that would ensure  
 271 compliance of a WQ time series realization. If  $c_p$  is below or equal to  $C^*$  then the WQ time series  
 272 is already compliant; if the  $c_p$  is above  $C^*$  then a concentration reduction of  $c_p - C^*$  is required  
 273 for the WQ time series to be compliant. (Required concentration reduction is  $C^* - c_p$  for dissolved  
 274 oxygen). Note that the  $G$  is estimated for each of the WQ concentration time series as if the  
 275 concentration values at different time-steps are independent draws of random variable with the  
 276 distribution function  $G$ . The  $M$  realizations of reconstructed WQ time series will yield  $M$  values  
 277 of  $c_p$  and, in turn, a distribution of  $c_p$ . Subsequently, Eq. (12) is used to compute the probability  
 278 of compliance. The probability of non-compliance ( $\beta_p$ ) for a given fraction of permissible  
 279 violations ( $p$ ) is defined as

$$\beta_p = 1 - \kappa_p. \quad (14)$$

280 Figure 2 shows a graphical illustration of computation of  $\kappa_p$  for a given value of concentration/load  
 281 reduction. Fig 2a shows the ensemble of distribution functions  $G$  of the 10000 reconstructed TP  
 282 concentration time series without any concentration/load reduction. At  $p = 0.05$ , the value of  $c_p$   
 283 obtained for one of the realizations of TP time series is shown. Fig 2b shows the histogram of  $c_p$   
 284 values obtained in this manner from the ensemble of reconstructed WQ time series. Clearly, all  
 285 the  $c_p$  values are above  $0.08 \text{ mg L}^{-1}$ ; therefore,  $\kappa_p$  is zero. Fig. 2c shows the histogram of  $c_p$   
 286 values after concentration reduction; the dark green area corresponds to the fraction of the  
 287 ensemble time series that violates the TMDL criterion  $0.08 \text{ mg L}^{-1}$  more than 0.05 fraction of of  
 288 the times after the concentration reduction. The DST lists percentage compliance ( $\kappa$ ) with different

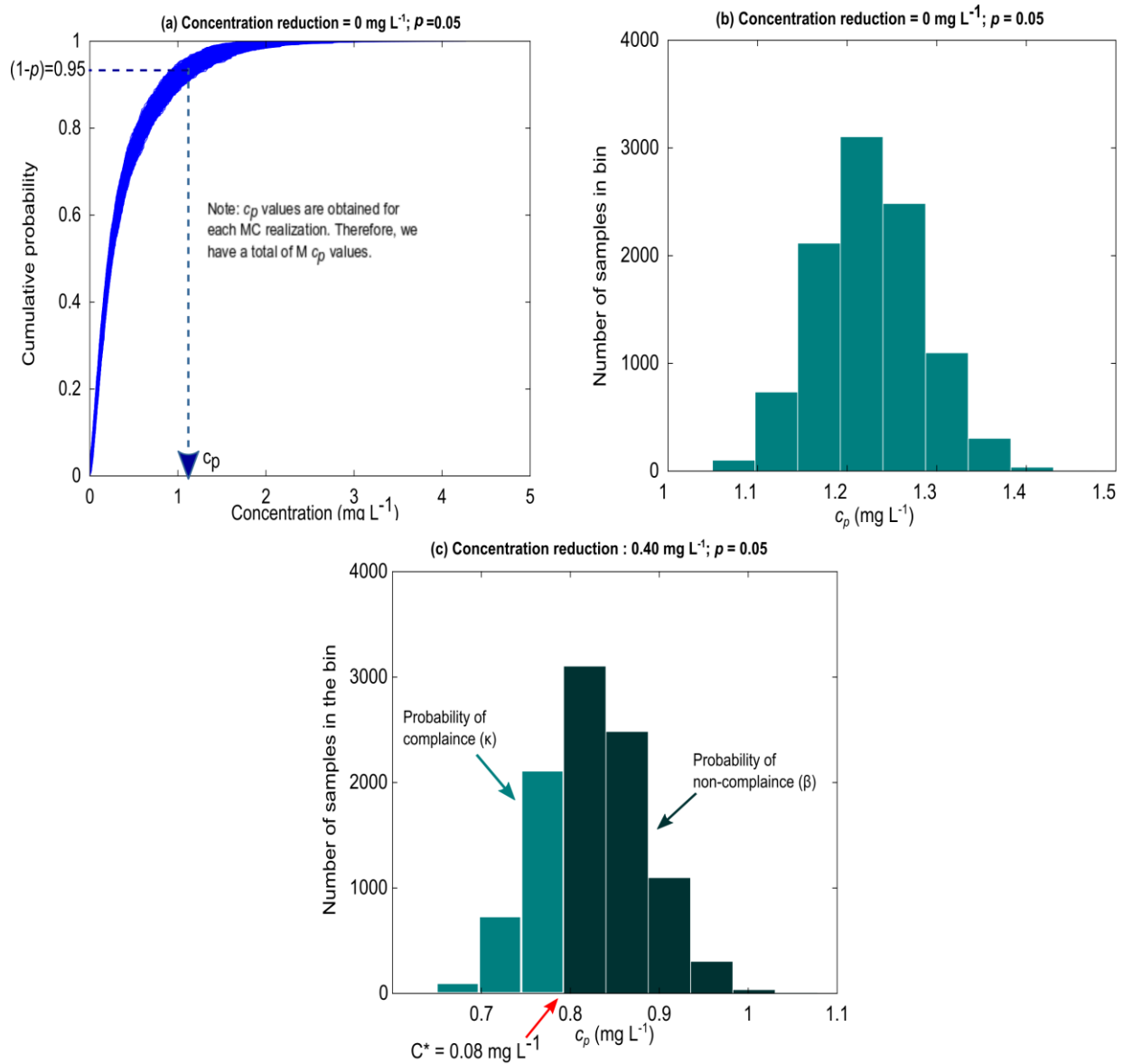
289 values of permissible violations ( $p$ ) at different load and concentration reduction in two tables.  
290 Hereafter, we drop the subscript  $p$  from  $\kappa$  and  $\beta$  for brevity.

291 The DST computes load reduction required (LRR) at a daily timescale. If the waterbody is required  
292 to be compliant at monthly or annual timescales, then LRR should be computed at monthly or  
293 annual timescales, respectively, and the WQ and TMDL time series should be aggregated from  
294 daily to monthly or annual timescales. Another way of computing LRR is to compute the difference  
295 between *average daily load* and *average TMDL load*. Average daily load is the average of the load  
296 over the entire time-period of analysis, and, similarly, the average TMDL is the average of the  
297 TMDL over the entire the period of analysis. The DST also computes average pollutant load in the  
298 waterbody and the difference between average daily load and average TMDL load. We note that  
299 the web based LOADEST tool computes the LRR by computing the difference between average  
300 simulated load and TMDL load but not the LRR at daily timescale.

301 A TMDL is established based on a threshold concentration below which a waterbody is no longer  
302 impaired. The consequences of water pollution are generally tied to concentration of pollutant in  
303 the water column rather than the total load carried by the waterbody. The value of concentration  
304 in a waterbody, however, depends upon the load introduced into it. Based on the application, either  
305 load and/or concentration reductions could be important. For example, a lake (especially a closed  
306 lake) ecosystem will be affected by both the load of TP in the lake-bed and concentrations of TP  
307 in the water column; but a river draining into the lake is likely to be affected only by concentration  
308 of the TP in the water column. In a river, high load of TP with high streamflow may result in low  
309 concentrations which will not affect the river ecosystem but could result in high load of TP to a  
310 receiving lake which will affect the lake ecosystem. The DST reports reductions required both in  
311 terms of constant concentration and constant load. Constant load reduction implies that measures  
312 are taken to reduce the load in the waterbody by the same amount at each day, and same for  
313 constant concentration reduction. However, reductions would be required only during periods  
314 when the waterbody violates the TMDL criterion. If nutrient violations are seasonal, then targeted  
315 pollution control measures only during these periods might achieve compliance.

316 The DST allows different values of the WQ standard (used in WQ risk analysis) and TMDL  
317 concentration (used in TMDL development). The distinction might be useful when WQ standard,  
318 usually determined by the federal agencies such as USEPA, cannot be met with available resources

319 so that a higher (in case of nutrients) or lower (in case of dissolved oxygen) TMDL concentration  
 320 value must be used.



321 Fig. 2. Illustration of risk-based total maximum daily load (TMDL) analysis for total phosphorus (TP) with target  
 322 TMDL concentration ( $C^*$ ) of  $0.08 \text{ mg L}^{-1}$ : (a) calculating  $c_p = F^{-1}(1 - p)$  for one of the realizations of  
 323 reconstructed water quality (WQ) series with no load or concentration reduction and  $p = 0.05$ , (b) histogram of  $c_p$   
 324 values before any load reduction, and (c) calculating the probability of compliance ( $\kappa$ ) from  $c_p$  values that were  
 325 obtained after a load reduction of  $210 \text{ kg day}^{-1}$  from the realizations of reconstructed WQ time series.

#### 326 2.4 What if streamflow data is not available at a water quality monitoring station?

327 Often, WQ monitoring stations in a river-network are not co-located with a streamflow gauge, but  
 328 a streamflow gauge may be available in proximity of a WQ monitoring station at a downstream or

329 upstream location in the river-network. In this case, the DST estimates streamflows,  $Q_u$ , at the  
330 ungauged station as (Emerson et al., 2005; Ries, 2007)

$$Q_u = \left( \frac{A_u}{A_g} \right)^b Q_g, \quad (15)$$

331 where  $A_u$  and  $A_g$  are the areas draining to the ungauged and gauged stations;  $Q_g$  is the measured  
332 flow rate at the streamflow gauge; and  $b$  is the exponent that varies with geographic region and  
333 climate but may be assumed to be equal to 1 when unknown. Emerson et al. (2005) reported the  
334 values of  $b$  equal to 0.85, 0.91, and 1.02 for winter, spring, and summer seasons, respectively, in  
335 Red River of the North Basin (in North Dakota and South Dakota). Typically,  $b = 1$  is accurate  
336 for mean-annual flows (Rodriguez-Iturbe and Rinaldo, 2001, chap. 1). The factor  $\left( \frac{A_u}{A_g} \right)^b$  in Eq.  
337 (15) is the watershed area ratio that must be supplied by the user. When streamflow measurements  
338 are available at the WQ sampling site, this conversion factor will be equal to unity. When the  
339 sampling site is located upstream of a streamflow gauge, the conversion factor will be less than  
340 unity; when the sampling site is located downstream of a streamflow gauge, the conversion factor  
341 will be greater than unity (Emerson et al., 2005; Ries, 2007). The watershed area ratio method  
342 assumes that ratio of total volume of water flowing through the ungauged station and gauged  
343 station is a function of drainage area of the two stations. At daily timescale, this assumption will  
344 be valid only if the drainage areas of the gauged and ungauged monitoring stations have *significant*  
345 overlap; that is, the streamwise distance between the two stations is small. This assumption is  
346 reasonable at annual timescale (Rodriguez-Iturbe and Rinaldo, 2001) but, at daily timescale, it is  
347 valid only in limited situations. At daily timescale, factors such as spatial variability of rainfall  
348 (Gabellani et al., 2007) and difference in land-use and topography will result in differences in the  
349 shape of streamflow time series. At daily timescale, a hydrologic model must be employed to  
350 estimate streamflows at ungauged stations. Gupta and Govindaraju (2019) showed that a simple  
351 hydrologic model calibrated against observations at a gauged location may entail significant  
352 uncertainties in estimated streamflows at ungauged locations. Nevertheless, the DST does not  
353 account for these uncertainties; this topic would be the scope of a different study. The DST  
354 operates at daily timescale; thus, streamflow and WQ data should be available at daily timescale.  
355 Typically, WQ data are collected as grab samples representing instantaneous values. An implicit  
356 assumption in the DST is that instantaneous values represent the average daily values.



### 357 3. Case study

#### 358 3.1 Study area and data

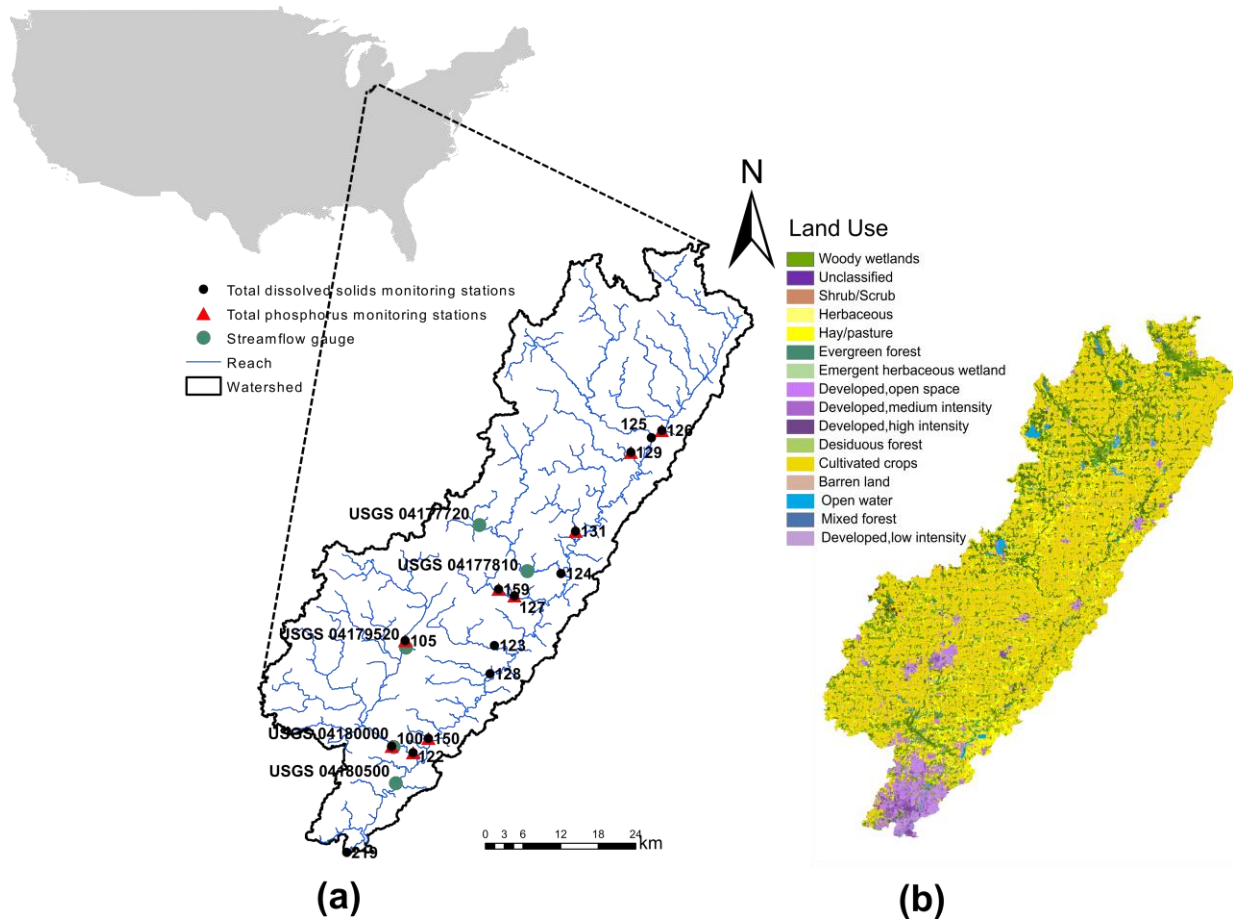
359 The DST was used to conduct WQ risk and TMDL analyses in the St. Joseph River Watershed  
360 (SJRW), USA (Fig. 3). The watershed is spread over parts of Indiana, Ohio, and Michigan. A large  
361 part of SJRW is covered with agricultural fields (Fig. 3b). Therefore, agricultural runoff is  
362 expected to be the main contributor of total phosphorus (TP) in the SJRW river-network (IDEM,  
363 2017, pp. 34). A few animal operations and point-sources, in urban areas and large villages, can  
364 potentially contribute TP in some parts of SJRW river-network. Total dissolved solids (TDS) are  
365 solid particles suspended in water column which may consist of nutrients, biological particles such  
366 as algae, and soil particles. Soil particles end up in water column because of channel erosion under  
367 high flow conditions, soil erosion from agricultural areas (livestock grazing, plowing), and urban  
368 areas (construction sites) and, to a limited extent, from forests (IDEM, 2017).

369 The daily-streamflow data at five stations in the watershed were made available by United States  
370 Geological Survey (USGS, 2016). The TP concentration data at 11 locations and TDS at 14  
371 locations in the watershed were made available through St. Joseph River Watershed Initiative  
372 (SJRWI at <http://wqis.ipfw.edu/>, accessed: 26 Aug, 2018, Figs. 4 and 5). St. Joseph River (SJR)  
373 and its tributaries are designated for aquatic life use (ALU), recreational use (RU), and warm and  
374 cold habitats (IDEM, 2017). TP and TDS primarily affect ALU; IDEM (2017) reported that many  
375 portions of the SJR and its tributaries were impaired for ALU. They concluded that TP and TDS  
376 load reductions of up to 66% and 95%, respectively, would be required to meet the TMDL  
377 criterion. Figs. 4 and 5 show that all the stations violated TP concentration standards except  
378 stations 123, 128, and 150, and all the stations violated TDS standards.

379 In this study, the WQ risk and TMDL analyses were restricted to time-period 2000-2017, thus TP  
380 and TDS data were reconstructed using observed streamflow from 01/01/2000 to 12/31/2017. It  
381 should be noted that (a) selection of this time-period is user's choice as long as there is a time-  
382 period where both streamflow and WQ have concurrent measurements. For each TP and TDS  
383 measurement station, streamflow data available at the nearest streamflow measurement station  
384 were used as an input to the DST. For example, for station 122, streamflow data available at the  
385 station USGS 04180000 were used. The required area-ratios (section 2.4) were computed using

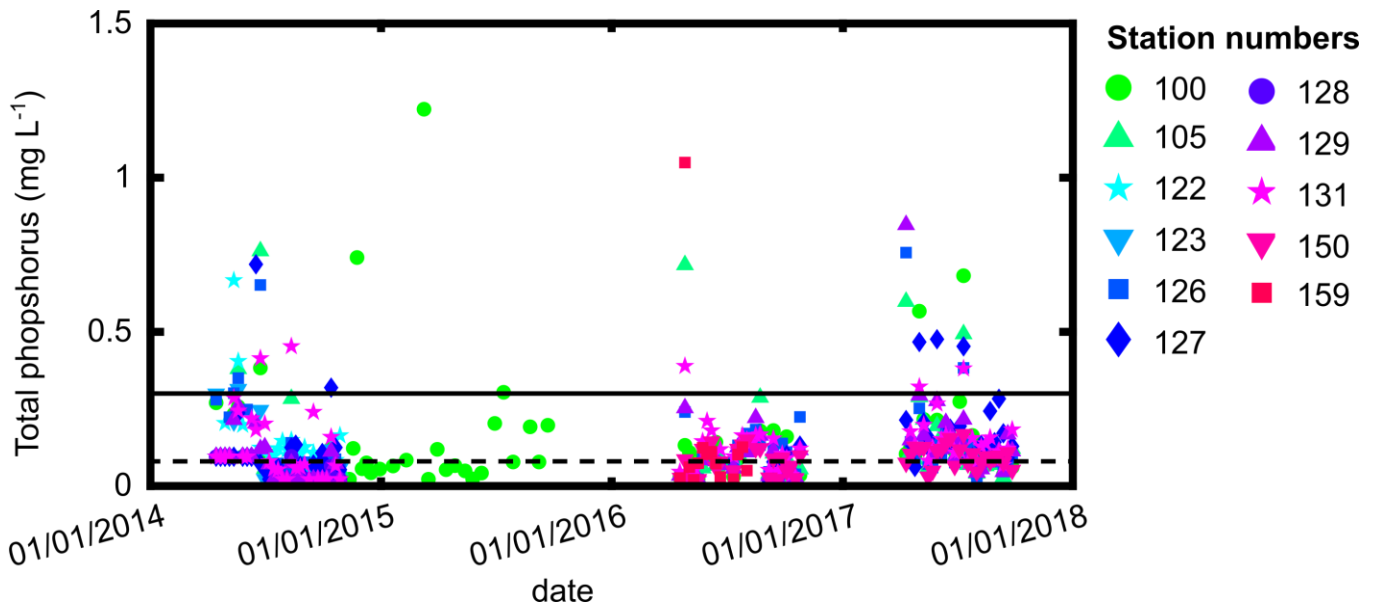
386 SWAT, though this ratio could be computed using any geographical information system or  
387 available drainage area information.

388 Concentrations values of  $0.08 \text{ mg L}^{-1}$  and  $750 \text{ mg L}^{-1}$  were used as WQ standards for calculation  
389 of WQ risk measures with respect to TP and TDS, respectively (SJRWI by <http://wqis.pfw.edu/>,  
390 accessed: 26 Aug, 2018); and concentration values of  $0.30 \text{ mg L}^{-1}$  and  $30 \text{ mg L}^{-1}$  were used for  
391 TMDL development for TP and TDS (IDEM, 2017), respectively. Out of the 9 LOADEST  
392 equations, the one with the best fit based on AIC was used to reconstruct WQ data. Subsequently,  
393 results corresponding to station 122 are discussed in detail and results corresponding to other  
394 stations are summarized for brevity. Henceforth, the analysis conducted using the DST described  
395 in this paper is referred to as rvm-LOADEST, and the analysis without uncertainty quantification  
396 is referred to as deterministic-LOADEST. The deterministic analysis was conducted using the  
397 web-based tool (by Park et al., 2015). The same LOADEST equation was used for reconstruction  
398 in both rvm- and deterministic-LOADEST. LOADEST equations used for TP and TDS  
399 reconstruction at different stations are listed in Appendix A. Note that the DST may also be used  
400 in a deterministic mode by using the expected value of the reconstructed WQ concentration/load  
401 time series and ignoring the information on uncertainty.

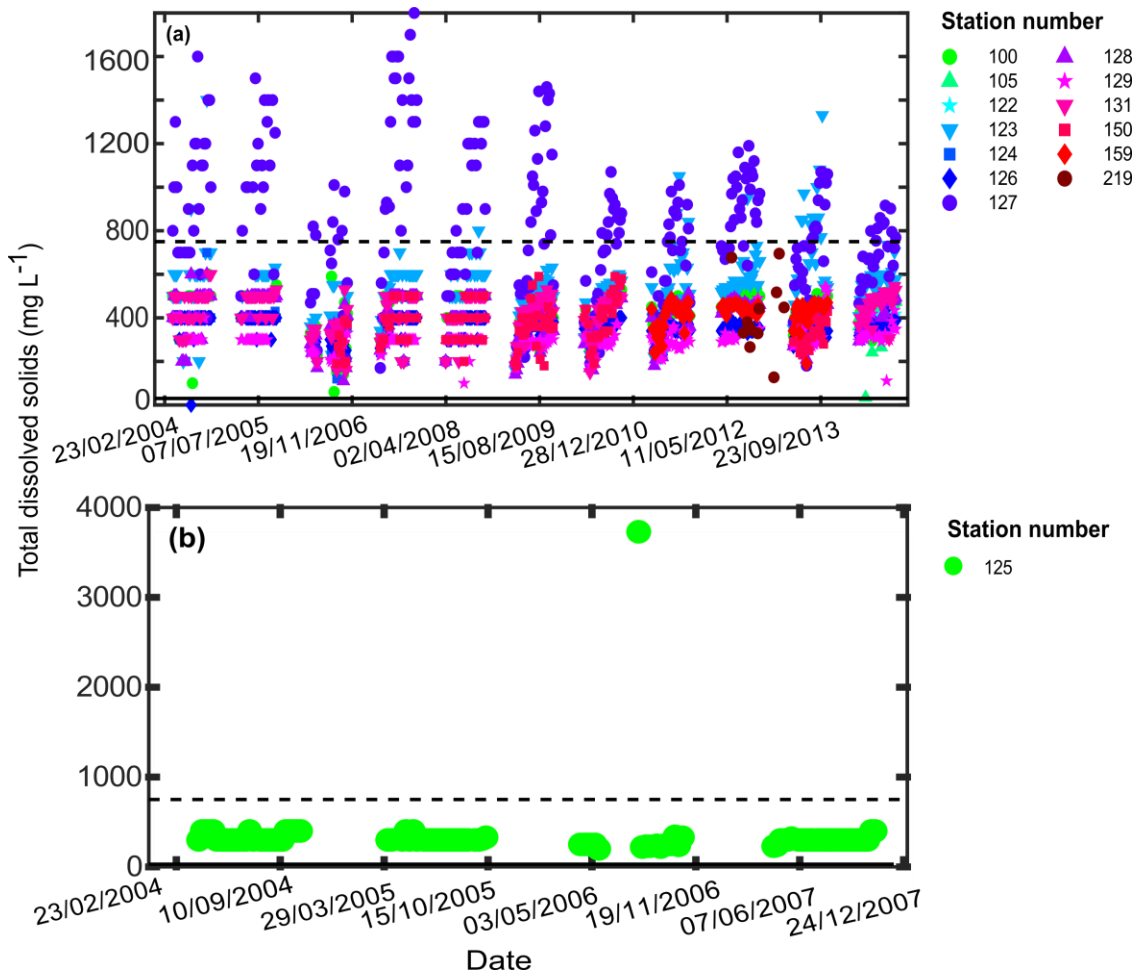


402 Fig. 3. (a) St. Joseph River watershed (SJRW) with a delineated stream network and locations of total phosphorus  
 403 (TP), total dissolved solids (TDS), and streamflow data stations. For TP and TDS, station numbers as listed on  
 404 SJRWI website are shown by red triangles and black dots, respectively, and for streamflow, USGS station numbers  
 405 are shown by green colored dots. (b) Land use pattern of SJRW.

406



407 Fig. 4. Observed total phosphorus data at 11 monitoring stations. The solid black line represents TMDL  
 408 concentration and dashed black line represents standard concentration.



409 Fig. 5. Observed total dissolved solids (TDS) data at (a) all the monitoring stations except 125 (b) monitoring station  
410 125. The observations at station 125 are shown separately because of one instance of exceptionally high value of  
411 TDS at this station. The solid black line represents TMDL concentration and dashed black line represents standard  
412 concentration.

### 413 3.2 Results

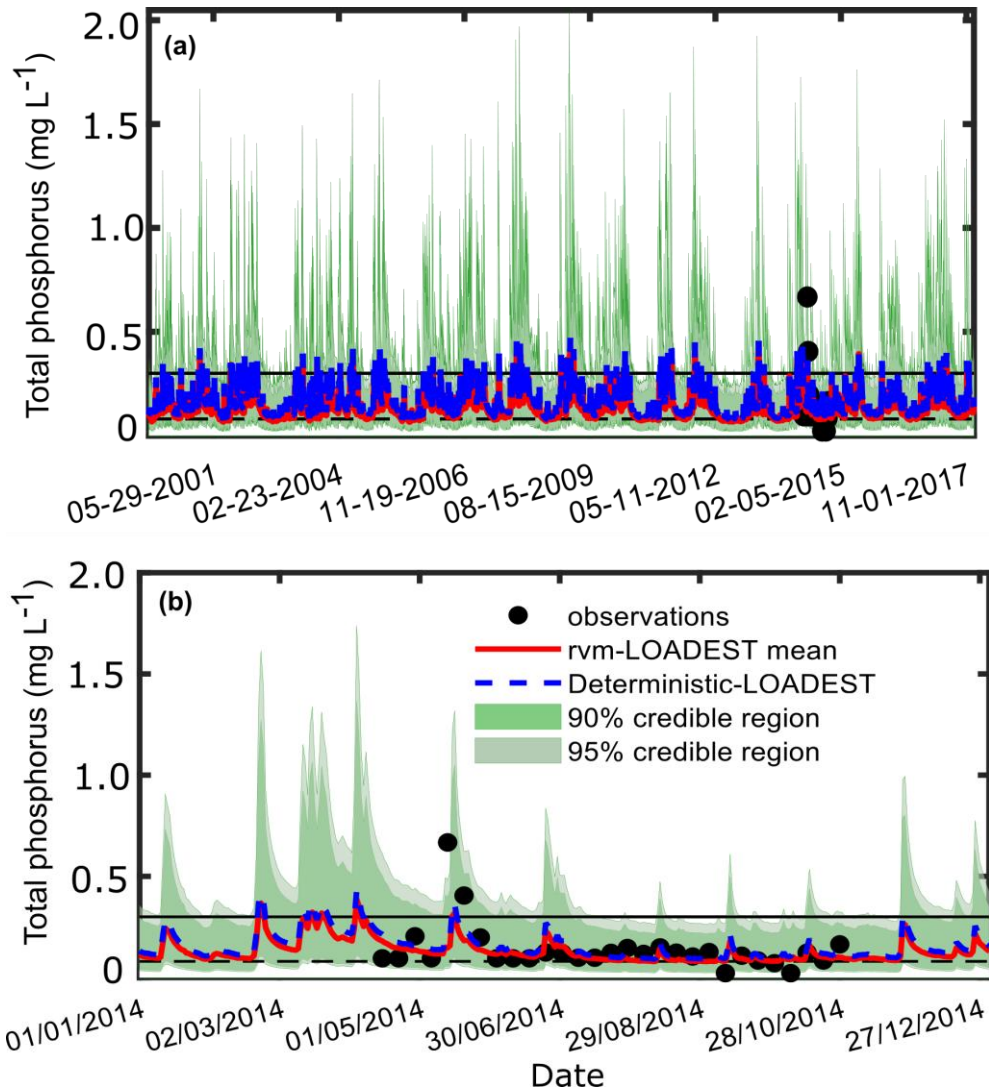
414 The run time of DST depends upon the time-period of analysis and number of observations  
415 available for reconstruction. The total runtime of the DST for one WQ monitoring station and one  
416 WQ constituent was approximately 2 minutes for this case study.

#### 417 *Total Phosphorus (TP)*

418 Except three, all observations were enveloped by the 90% and 95% credible regions which is  
419 expected since all these observations were used for estimation of the weight vector and variance  
420 of residuals (Fig. 6). The uncertainty band was wide, especially in high concentration regions.  
421 Since the land-use in SJRW is dominated by agriculture, high concentrations of TP are expected  
422 to be associated with high agricultural runoff and high streamflows. According to Eq. (4), high  
423 streamflows will result in high prediction variance of TP concentrations. The reconstructed TP  
424 time series obtained by the computing mean of rvm-LOADEST time-series and obtained by  
425 deterministic-LOADEST were approximately the same (Fig. 6). The estimated weights by RVM  
426 were consistently smaller than those obtained by LOADEST, providing better hedging against  
427 errors in streamflow observations (Appendix A1).

428  
429 The ranges of R-R-V and WH measures at station 122 were 0.245 to 0.278 (reliability), 0.235 to  
430 0.278 (resilience), 0.572 to 0.593 (vulnerability) and 0.289 to 0.320 (watershed health),  
431 respectively (Table 2). The R-R-V and WH measures obtained by deterministic-LOADEST were  
432 0.007, 0.002, 0.468 and 0.020, respectively. The mean values of the measures obtained by rvm-  
433 LOADEST are substantially different from those obtained by deterministic-LOADEST. In fact,  
434 the values of the risk measures obtained by deterministic-LOADEST are not even contained in the  
435 range of those obtained by rvm-LOADEST; this is due to consideration of uncertainty in rvm-  
436 LOADEST and the sensitivity of the nonlinear risk measures to slight differences in the WQ  
437 constituent time series. At station 122, the reliability and resilience are low, and the vulnerability  
438 is high. It implies that this station incurs frequent violations of WQ standard and takes a long time  
439 to recover, and severity of violations is also high. Note that if pollution control measures are used

440 to increase the resilience of station 122, the reliability would also increase. However, if pollution  
441 control measures are taken to increase the reliability only, the resilience may not increase.  
442 Figure 7 shows the LRR computed by deterministic- and by rvm-LOADEST at different  
443 compliance values and different permissible violations. At most of the stations, the LRR as  
444 computed by using deterministic-LOADEST did not achieve even 50% compliance after the  
445 uncertainty in reconstructed WQ time series was considered. For example, at station 122, the LRR  
446 as obtained by deterministic- and rvm-LOADEST at  $p = 0.10$  and 50% compliance were 0 and  
447  $35.9 \text{ Kg day}^{-1}$ , respectively; the LRR as obtained by deterministic- and rvm-LOADEST at  $p =$   
448  $0.05$  and 50% compliance were 0 and  $211.6 \text{ Kg day}^{-1}$ , respectively; the LRR as obtained by  
449 deterministic- and rvm-LOADEST at  $p = 0.03$  and 50% compliance were  $75.9 \text{ Kg day}^{-1}$  and  
450  $458.5 \text{ Kg day}^{-1}$ , respectively. The LRR values as obtained by rvm-LOADESTs are function of  $\kappa$ ;  
451 as the  $\kappa$  increases, the LRR increases.  
452  
453



454 Fig. 6. Station 122, total phosphorus (TP). Observed and reconstructed daily TP during (a) 2000-2017 and (b) 2014  
 455 (the year in which water quality data was available)

456

457

458 Table 2. Total phosphorus (TP). The risk-measures obtained by deterministic (D)- and rvm-LOADEST. The  
 459 measures are computed at daily timescale.

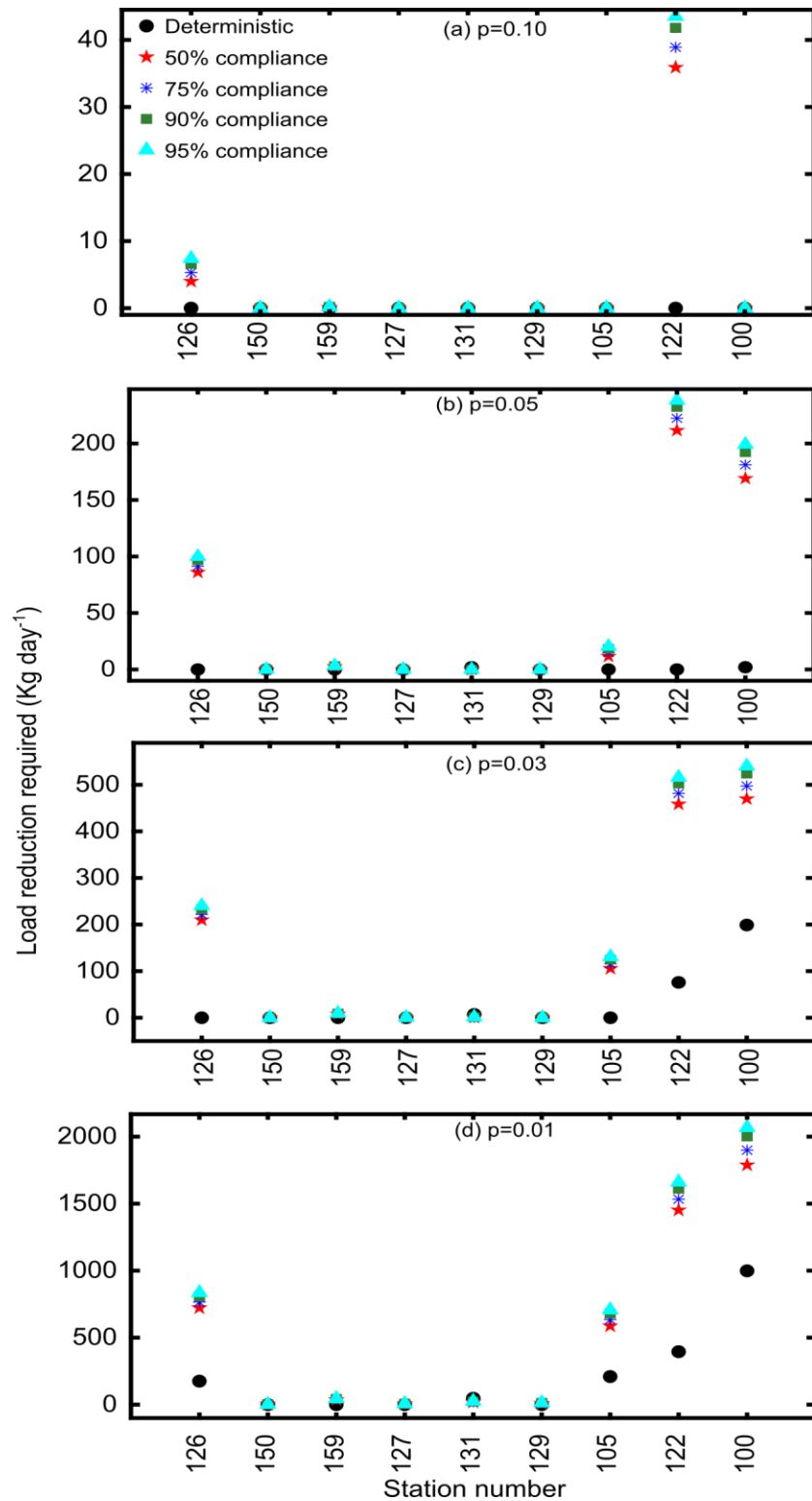
Station	Model Type	RVM statistic	Reliability	Resilience	Vulnerability	Watershed health
126	D	-	0.183	0.023	0.424	0.134
	rvm	(Mean, Median)	(0.363, 0.363)	(0.342, 0.342)	(0.557, 0.557)	(0.381, 0.381)
		Range	(0.341-0.382)	(0.316-0.368)	(0.545-0.568)	(0.364-0.396)
150	D	-	0.962	0.008	0.202	0.182
	rvm	(Mean, Median)	(0.787, 0.787)	(0.660, 0.660)	(0.448, 0.448)	(0.659, 0.659)
		Range	(0.769-0.803)	(0.614-0.705)	(0.417-0.472)	(0.637-0.682)
159	D	-	0.001	0.001	0.300	0.009
	rvm	(Mean, Median)	(0.433, 0.433)	(0.433, 0.433)	(0.587, 0.587)	(0.426, 0.426)
		Range	(0.410-0.455)	(0.404-0.460)	(0.575-0.598)	(0.407-0.442)
127	D	-	0.720	0.034	0.412	0.244

	<b>rvm</b>	(Mean, Median) Range	(0.642, 0.642) (0.622-0.659)	(0.453, 0.453) (0.418-0.490)	(0.482, 0.482) (0.465-0.498)	(0.532,0.532) (0.515-0.551)
131	<b>D</b>	-	0.428	0.030	0.500	0.180
	<b>rvm</b>	(Mean, Median) Range	(0.596, 0.596) (0.578-0.617)	(0.465, 0.465) (0.431-0.498)	(0.470, 0.470) (0.454-0.484)	(0.528, 0.528) (0.510-0.546)
129	<b>D</b>	-	0.813	0.05 0	0.373	0.29
	<b>rvm</b>	(Mean, Median) Range	(0.776, 0.776) (0.760-0.791)	(0.478, 0.478) (0.428-0.529)	(0.448, 0.448) (0.428-0.473)	(0.589, 0.589) (0.564-0.615)
105	<b>D</b>	-	0. 807	0.132	0.476	0.382
	<b>rvm</b>	(Mean, Median) Range	(0.658, 0.658) (0.633-0.680)	(0.493, 0.493) (0.446-0.537)	(0.521, 0.521) (0.502-0.541)	(0.537, 0.538) (0.516-0.560)
122	<b>D</b>	-	0.007	0.002	0.47	0.020
	<b>rvm</b>	(Mean, Median) Range	(0.265, 0.266) (0.245-0.285)	(0.256, 0.256) (0.235-0.278)	(0.582, 0.582) (0.572-0.593)	(0.305, 0.305) (0.289-0.320)
100	<b>D</b>	-	0.464	0.062	0.451	0.250
	<b>rvm</b>	(Mean, Median) Range	(0.477, 0.477) (0.457-0.496)	(0.398, 0.399) (0.370-0.431)	(0.542, 0.543) (0.530-0.556)	(0.443, 0.443) (0.426-0.458)

460

461





462 Fig. 7. Total phosphorus (TP). Daily timescale load reduction required at different station at  
 463 permissible violations ( $p$ )  
 464

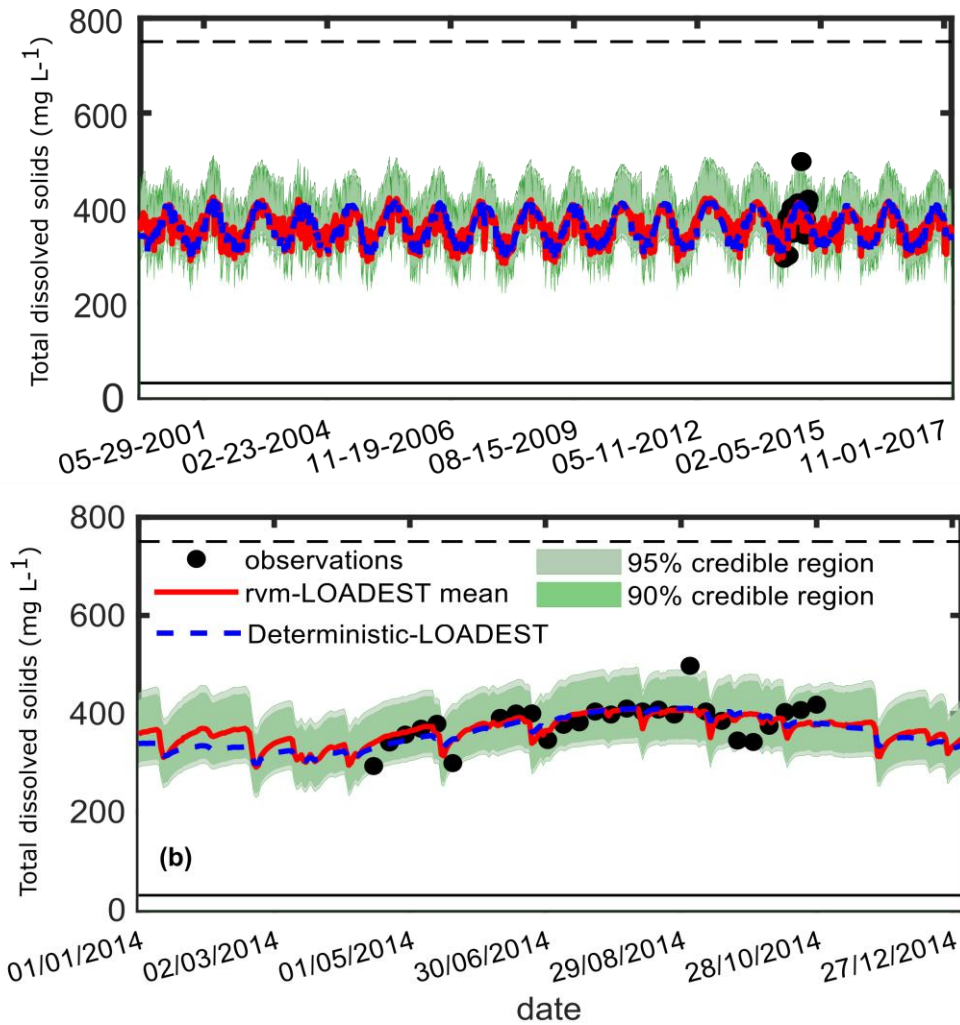
465 *Total Dissolved Solids (TDS)*

466 The general results obtained for TDS were same as those for TP. The reconstructed TDS time  
467 series obtained by mean of the rvm-LOADEST and deterministic-LOADEST were similar at most  
468 time-steps (Fig. 8b). All except one observation were enveloped by the 90% and 95% credible  
469 regions. Table 3 lists the statistics of R-R-V values obtained by using deterministic- and rvm-  
470 LOADEST models. As in case of TP, at many stations, the R-R-V values obtained by  
471 deterministic-LOADEST were outside the range of those obtained by rvm-LOADEST. The  
472 differences between the risk measured obtained by deterministic-LOADEST and the means of the  
473 risk measures obtained by rvm-LOADEST, however, was small. TDS violations occurred only at  
474 a few stations in the watershed. When violations did occur, the resilience of the station was low.  
475 For example, at station 127, the reliability was 0.604 (mean of rvm-LOADEST values) and the  
476 resilience was 0.353 implying if pollution control measures were to be put in place to increase the  
477 resilience of the waterbody, these will also increase the reliability of the waterbody. At most  
478 stations, the vulnerability was low implying that magnitude of violations was small. Overall, the  
479 watershed is in a healthy condition in terms of TDS.

480 At stations 105 and 122, the deterministic-LOADEST LRR values were larger than rvm-  
481 LOADEST LRR values computed at 50% and higher levels of compliance. At most of the stations,  
482 however, the LRR values computed by deterministic-LOADEST at various compliance  
483 percentages were smaller than those computed by deterministic-LOADEST (Table 5). In  
484 summary, if one carries out a deterministic analysis then the computed LRR may not be enough to  
485 comply with TMDL concentration, or may be too conservative as is the case for stations 105 and  
486 122.

487

488



489 Fig. 8. Station 122, total dissolved solids (TDS). Observed and reconstructed daily TDS during (a) 2000-2017 and  
 490 (b) 2014 (the year in which water quality data were available)

491

492 Table 3. Total dissolved solids (TDS). The risk-measures obtained by deterministic (D)- and rvm-LOADEST. The  
 493 risk measures are computed at daily timescale.

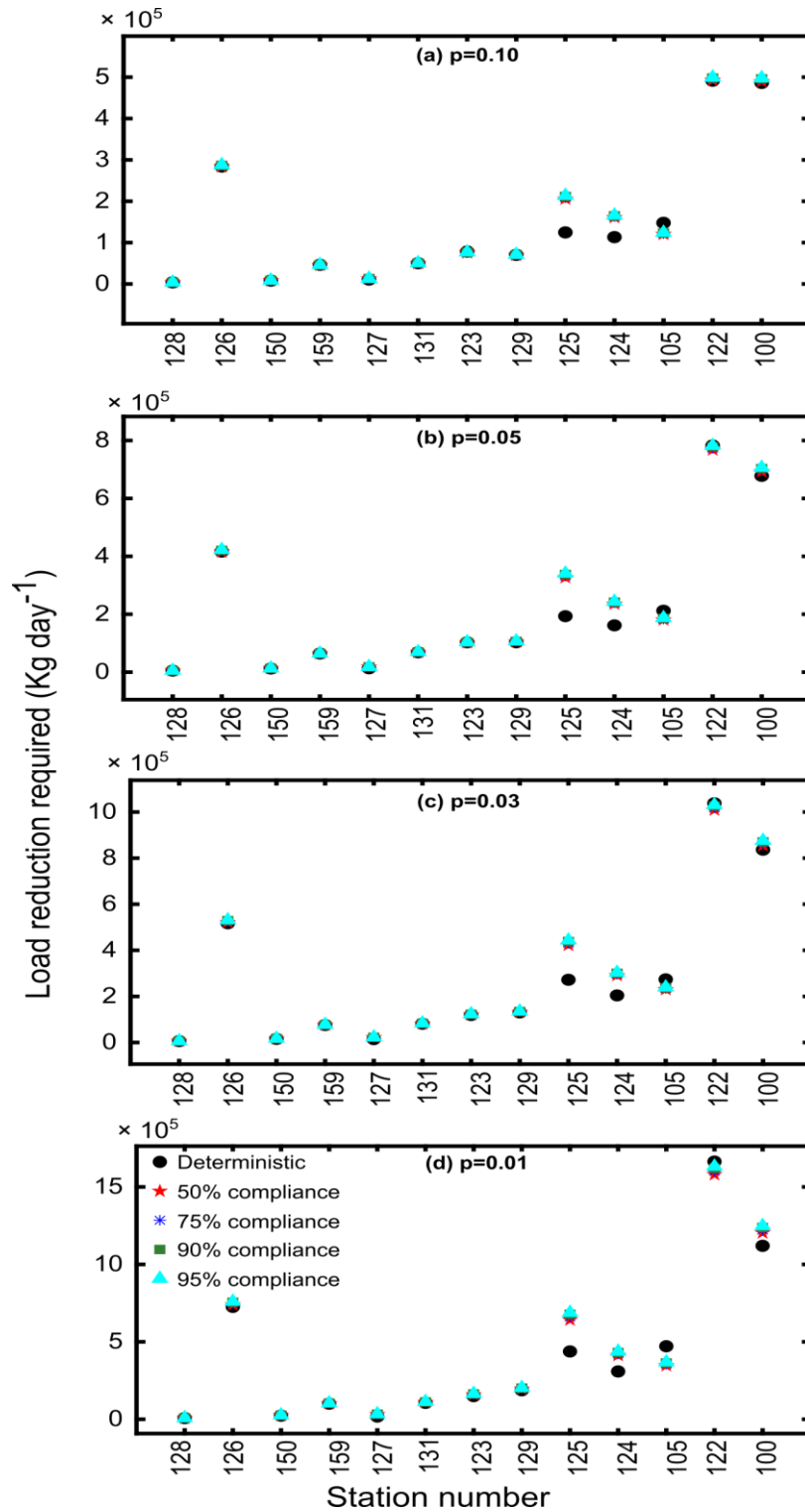
Station	Model type	RVM statistic	Reliability	Resilience	Vulnerability	Watershed health
128	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median)	(1.000, 1.000)	(0.999, 1.000)	(0.004, 0.000)	(0.999, 1.000)
		Range	(1.000)	(0.500-1.000)	(0.000-0.135)	(0.786-1.000)
126	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median)	(1.000, 1.000)	(1.000, 1.000)	(0.000, 0.000)	((1.000, 1.000))
		Range	(1.000)	(1.000)	(0.000)	(1.000)
150	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median)	(0.999, 1.000)	(0.999, 1.000)	(0.015, 0.000)	(0.995, 1.000)
		Range	(0.999-1.000)	(0.500-1.000)	(0.000-0.221)	(0.769-1.000)
159	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median)	(1.000, 1.000)	(1.000, 1.000)	(0.000, 0.000)	(1.000, 1.000)
		Range	(1.000)	(1.000)	(0.000)	(1.000)
127	<b>D</b>	-	0.581	0.053	0.210	0.29

	<b>rvm</b>	(Mean, Median) Range	(0.604, 0.605) (0.588-0.623)	(0.354, 0.354) (0.324-0.390)	(0.298, 0.298) (0.286-0.311)	(0.532, 0.532) (0.514-0.552)
131	<b>D</b>	-	0.996	0.115	0.020	0.483
	<b>rvm</b>	(Mean, Median) Range	(0.986, 0.986) (0.981-0.991)	(0.860, 0.861) (0.696-1.000)	(0.066, 0.066) (0.045-0.089)	(0.925, 0.926) (0.859-0.976)
123	<b>D</b>	-	0.957	0.094	0.055	0.44
	<b>rvm</b>	(Mean, Median) Range	(0.898, 0.899) (0.884-0.911)	(0.713, 0.713) (0.650-0.779)	(0.107, 0.107) (0.097-0.122)	(0.830, 0.830) (0.802-0.857)
129	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median) Range	(0.999, 0.999) (0.998-1.000)	(0.965, 1.000) (0.333-1.000)	(0.0497, 0.048) (0.000-0.229)	(0.971-0.982) (0.680-1.000)
125	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median) Range	(1.000, 1.000) (1.000)	(1.000, 1.000) (1.000)	(0.000, 0.000) (0.000-0.058)	(0.1000, 1.000) (0.980-1.000)
124	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median) Range	(0.999, 1.000) (0.996-1.000)	(0.998, 1.000) (0.50-1.00)	(0.034, 0.029) (0.000-0.240)	(0.987, 0.990) (0.764-1.000)
105	<b>D</b>	-	0.999	0.125	0.054	0.491
	<b>rvm</b>	(Mean, Median) Range	(0.915, 0.915) (0.903-0.928)	(0.769, 0.769) (0.693-0.835)	(0.165, 0.165) (0.146-0.184)	(0.837, 0.839) (0.807-0.864)
122	<b>D</b>	-	1.000	1.000	0.000	1.000
	<b>rvm</b>	(Mean, Median) Range	1.000 (1.000)	1.000 (1.000)	0.000 (0.000)	1.000 (1.000)
100	<b>D</b>	-	1.000	1.000	0.000	1.00
	<b>rvm</b>	(Mean, Median) Range	(0.996, 0.998) (0.995-1.000)	(0.981, 1.000) (0.750-1.000)	(0.045, 0.044) (0.014-0.122)	(0.978, 0.983) (0.898-0.995)

Output not available from web-based LOADEST for station 219

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495



496 Fig. 9. Total dissolved solids (TDS). Daily time scale load reduction required at different station  
 497 at permissible violations ( $p$ )

498 3.3 Discussion

499 In case of TP, the R-R-V measures obtained by deterministic-LOADEST were not even in the  
500 range of those obtained by rvm-LOADEST. This discrepancy is due to the assumption of statistical  
501 independence of errors in the reconstructed TP series at different time-steps and high sensitivity  
502 of these measures when the WQ concentration series is near the standard series. The discrepancy  
503 exists even though the reconstructed TP time series obtained by deterministic-LOADEST was  
504 approximately equal to the mean of the TP time series obtained by rvm-LOADEST. Out of 10,000  
505 MC realizations of TP time series, there were many realizations that were close to the mean  
506 realization. But no realization was identical to the mean time series because of the added error-  
507 term in rvm-LOADEST. The difference between mean and actual realization of TP time series  
508 translated into large differences in estimated R-R-V measures because these measures are  
509 extremely sensitive to small changes in TP time series, especially when TP time series is close to  
510 the standard time series. Thus, if one were to simply take the mean values estimated by LOADEST  
511 and not account for uncertainty in these estimates, the assessment of WQ risk may be erroneous.  
512 Since the discrepancy in R-R-V measures obtained by rvm-LOADEST and deterministic-  
513 LOADEST is due to statistical assumptions over the residual time series, this discrepancy also  
514 illustrates the importance of the assumptions in uncertainty analysis. This discrepancy can be  
515 resolved only if the statistical assumptions are such that they allow to draw the mean WQ  
516 constituent time series from the distribution. Moreover, evidence supports the hypothesis that a  
517 watershed that yields a WQ constituent time series with high positive autocorrelation is less  
518 resilient to perturbations (such as high pollutant load in the watershed) (Qi et al., 2016). Thus,  
519 introducing artificial correlations to reconstructed WQ constituent time series by means of  
520 statistical assumptions over the error-term may lead to misleading conclusions. There is no way to  
521 confirm if the statistical assumptions made are valid; QQ plots can tell us if the assumption are  
522 invalid, not if the assumptions are correct. One of the problems in probabilistic uncertainty  
523 quantification is the non-uniqueness of possible statistical assumptions that can fit the data.  
524 To check the effectiveness of MOS value used in TMDL reports (e.g., IDEM, 2017), the MOS  
525 was computed using the same method as in IDEM (2017) (see SI). Specifically, five different  
526 TMDLs were computed for five different flow regions in each of the streams. The five flow regions  
527 were determined using daily exceedance probability (DEP) as follows : high flows (0 – 0.10  
528 DEP), moist conditions (0.10 – 0.40 DEP), mid-range (0.40 – 0.60 DEP), dry conditions (0.60-  
529 0.90 DEP), and low flows (0.90 – 1.00 DEP). The results are discussed for high flow conditions

530 only. As per rvm-LOADEST, up to 1000% of the TMDL should be allocated as MOS to achieve  
531 50% compliance (Figs. S3 and S4). An MOS value greater than 100% implies that the maximum  
532 allowable load in the watershed cannot be determined reliably. IDEM (2017) report did not  
533 explicitly quantify uncertainties. The cases study illustrates how a deterministically estimated WQ  
534 times-series may lead to different conclusions than one that quantifies the effects of uncertainty.  
535 Authors of IDEM (2017) report used SWAT to simulate WQ constituent time series. SWAT is  
536 known to incur high uncertainties (Hollaway et al., 2018) but this information was not utilized in  
537 MOS specification.

538 To reconstruct WQ constituent time series, the DST assumes that logarithm of load and logarithm  
539 of streamflow are correlated. If the correlation is weak, the prediction accuracy of the model will  
540 be poor and the uncertainty band will be wide. To check the prediction accuracy of rvm-  
541 LOADEST, the observed and predicted mean values of TP and TDS were plotted; the results  
542 suggested that one can indeed reconstruct TP and TDS loads by using streamflow values as  
543 predictor variables (Figs. S5 and S6). In case of both TP and TDS, 90 to 100% of the observations  
544 were enveloped by the 90% credible region with a few exceptions (Tables S3 and S4). Moreover,  
545 the methodology adopted in the DST is valid only if the relationship between streamflows and  
546 pollutant loads remains unchanged during the period of analysis. If this relationship changes,  
547 model predictions will be poor. The relationship could change because of pollution control  
548 measures put in the drainage area during the period of analysis, but not during the observation  
549 period and change of rainfall-runoff-pollutant load relationship (due to climatic and/or land use  
550 changes), Thus, the DST cannot be used for future scenario analyses indiscriminately.

551 Further, the DST assumes that the residuals at different time-steps are statistically independent. If  
552 this assumption is invalid, the consequence would be over-estimation of information content in  
553 residuals which would result in under-estimation of uncertainty in  $w$  and, in-turn, an under-  
554 estimation of uncertainty in predicted loads. For convenience of analysis, the DST assumes that  
555 residuals are distributed according to Gaussian law with homoscedastic variance. To check the  
556 validity of these statistical assumptions, QQ plots were used (Figs. S7-S10). These plots revealed  
557 that the observed residuals did not satisfy the assumptions made by DST, implying that  
558 uncertainties in WQ reconstruction as reported in this study may be underestimated. One way of  
559 relaxing the assumption of independence is to model the residuals as an autoregressive (AR)  
560 process (e.g., Hantush and Chaudhary, 2014). In this process, the residual at a time-step is

561 regressed against  $(k - 1)$  residuals at previous consecutive time-steps. Future versions of DST  
562 will be updated to accommodate this analysis.

563 Data limitations are ubiquitous in modeling exercises. In principle, there is no established  
564 restriction on the minimum number of observations to apply the DST. However, since the DST  
565 reconstructs WQ time series using a statistical regression method, very few observations may  
566 translate into over- or under-estimation of uncertainty. More observations imply better uncertainty  
567 estimates. For reference, Schwarz et al. (2006) suggested 15 observations to estimate *annual*  
568 *average* loads, and we suggest this number as a lower limit.

#### 569 **4. Summary and conclusions**

570 A DST was developed to reconstruct WQ constituent time series and conduct risk-based WQ  
571 assessment and TMDL development. The DST uses RVM to incorporate uncertainty in  
572 reconstruction of WQ constituent time series. The tool estimates uncertainty due to residual errors  
573 in reconstructed WQ time series, allows users to propagate this uncertainty to R-R-V and WH risk  
574 assessment and TMDL estimation. These two applications of the DST were demonstrated for the  
575 SJRW. The following conclusion were drawn:

576 (1) The weights estimated by RVM are consistently smaller than those compared to web based  
577 LOADEST; the smaller weights are desirable because they hedge against errors in  
578 streamflows.

579 (2) Based on our experience, we expect WQ risk measures to be very sensitive to small  
580 changes in WQ constituent time series especially when realizations of loads/concentrations  
581 are close to the standard values; therefore, errors in the reconstruction of WQ constituent  
582 time series must be modeled to obtain a realistic estimate of WQ risk measures of a  
583 waterbody. This sensitivity, however, also illustrates the importance of realistic statistical  
584 assumptions over the error-term.

585 (3) At most stations, consideration of uncertainty in WQ risk measures led to very different  
586 conclusions about watershed health. Uncertainty analysis indicated a relatively poorer  
587 health at some WQ monitoring stations and a relatively better health at other WQ  
588 monitoring stations.

589 (4) The LRR values at daily timescale as yielded by a deterministic analysis may not be enough  
590 to achieve even 50% compliance. Uncertainty in LRR should be considered for effective



591 pollution control. The arbitrarily selected MOS values used in TMDL report may result in  
592 gross under-estimation of uncertainty. Therefore, MOS values should be based upon a  
593 systematic uncertainty analysis (as was also suggested by Reckhow, 2003).

594 As presented, the tool is restricted to quantifying uncertainty due to residual errors. It is possible  
595 to use the RVM methodology to explicitly incorporate measurement errors in streamflows, which  
596 will be the topic of future research.

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## APPENDIX A

763 Table A1. List of LOADEST equations used for total phosphorus (TP) reconstruction at different  
764 monitoring stations in the study area

Station	TP	Weights	Variance of $\ln y$
100	4	D: $\ln y = 3.64 + 1.62\ln Q - 0.26 \sin(2\pi\delta t) - 0.30 \cos(2\pi\delta t)$ RVM: $\ln y = 3.67 + 1.58\ln Q - 0.17 \sin(2\pi\delta t) - 0.22 \cos(2\pi\delta t)$	D: 0.52 RVM: 0.52
105	8	D: $\ln y = 1.15 + 1.39\ln Q + 0.13\ln Q^2 - 0.28 \sin(2\pi\delta t) - 0.39 \cos(2\pi\delta t) + 0.12\delta t$ RVM: $\ln y = 1.30 + 1.35\ln Q + 0.11\ln Q^2 - 0.11 \sin(2\pi\delta t) - 0.24 \cos(2\pi\delta t) + 0.10\delta t$	D: 0.49 RVM: 0.49
122	1	D: $\ln y = 3.65 + 1.35\ln Q$ RVM: $\ln y = 3.65 + 1.33\ln Q$	D: 0.35 RVM: 0.34
126	1	D: $\ln y = 3.17 + 1.38\ln Q$ RVM: $\ln y = 3.16 + 1.37\ln Q$	D: 0.50 RVM: 0.50
127	3	D: $\ln y = -0.85 + 1.28\ln Q + 0.17\delta t$ RVM: $\ln y = -0.85 + 1.27\ln Q + 0.15\delta t$	D: 0.46 RVM: 0.45
129	3	D: $\ln y = 0.82 + 1.28\ln Q + 0.20\delta t$ RVM: $\ln y = 0.82 + 1.28\ln Q + 0.18\delta t$	D: 0.47 RVM: 0.47
131	7	D: $\ln y = 0.013 + 1.30\ln Q - 0.51 \sin(2\pi\delta t) - 0.61 \cos(2\pi\delta t) + 0.09\delta t$ RVM: $\ln y = 0 + 1.30\ln Q - 0.50 \sin(2\pi\delta t) - 0.62 \cos(2\pi\delta t) + 0.05\delta t$	D: 0.45 RVM: 0.44

150	7	D: $\ln y = -1.35 + 1.01\ln Q - 0.03\sin(2\pi\delta t) - 0.54\cos(2\pi\delta t) + 0.30\delta t$	D: 0.33
		RVM: $\ln y = -1.26 + 1.00\ln Q + 0\sin(2\pi\delta t) - 0.41\cos(2\pi\delta t) + 0.24\delta t$	RVM: 0.32
159	1	D: $\ln y = 0 + 1.02\ln Q$	D: 1.11
		RVM: $\ln y = 0.000014 + 1.02\ln Q$	RVM: 1.05

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Table A2. TP. Estimated covariance matrix of weights at station 122 using RVM

	$a_0$	$a_1$
$a_0$	0.013380	0.005026
$a_1$	0.005026	0.016743

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Table A3. List of LOADEST equations used for total dissolved solids (TDS) reconstruction at different monitoring stations in the study area

Station	TDS	Weights	Variance of $\ln y$
100	6	D: $\ln y = 12.04 + 0.82\ln Q - 0.02\ln Q^2 + 0.02\sin(2\pi\delta t) + 0.08\cos(2\pi\delta t)$ RVM: $\ln y = 12.05 + 0.83\ln Q - 0.02\ln Q^2 + 0.01\sin(2\pi\delta t) + 0.07\cos(2\pi\delta t)$	D: 0.034 RVM=0.034
105	1	D: $\ln y = 10.21 + 0.76\ln Q$ RVM: $\ln y = 10.21 + 0.75\ln Q$	D: 0.19 RVM: 0.19
122	4	D: $\ln y = 11.62 + 0.97\ln Q + 0.05\sin(2\pi\delta t) + 0.09\cos(2\pi\delta t)$ RVM: $\ln y = 11.65 + 0.94\ln Q + 0\sin(2\pi\delta t) + 0.04\cos(2\pi\delta t)$	D: 0.01 RVM: 0.01
123	9	D: $\ln y = 9.76 + 0.84\ln Q - 0.05\ln Q^2 - 0.04\sin(2\pi\delta t) + 0.03\cos(2\pi\delta t) + 0.01\delta t + 0\delta t^2$ RVM: $\ln y = 9.75 + 0.84\ln Q - 0.05\ln Q^2 - 0.02\sin(2\pi\delta t) + 0\cos(2\pi\delta t) + 0.01\delta t + 0\delta t^2$	D: 0.04 RVM: 0.04
124	3	D: $\ln y = 10.75 + 0.85\ln Q - 0.03\delta t$ RVM: $\ln y = 10.75 + 0.85\ln Q - 0.03\delta t$	D: 0.05 RVM: 0.05
125	1	D: $\ln y = 10.80 + 1.00\ln Q$ RVM: $\ln y = 10.80 + 1.00\ln Q$	D: 0.07 RVM: 0.07
126	9	D: $\ln y = 11.37 + 0.92\ln Q - 0.03\ln Q^2 + 0\sin(2\pi\delta t) + 0\cos(2\pi\delta t) - 0.01\delta t + 0\delta t^2$ RVM: $\ln y = 11.37 + 0.92\ln Q - 0.03\ln Q^2 + 0\sin(2\pi\delta t) + 0\cos(2\pi\delta t) - 0.01\delta t + 0\delta t^2$	D: 0.01 RVM: 0.01
127	7	D: $\ln y = 8.14 + 0.73\ln Q - 0.05\sin(2\pi\delta t) - 0.04\cos(2\pi\delta t) - 0.027\delta t$ RVM: $\ln y = 7.99 + 0.71\ln Q + 0\sin(2\pi\delta t) + 0\cos(2\pi\delta t) - 0.03\delta t$	D: 0.09 RVM: 0.10
128	2	D: $\ln y = 6.79 + 0.85\ln Q - 0.04\ln Q^2$ RVM: $\ln y = 6.80 + 0.85\ln Q - 0.04\ln Q^2$	D: 0.02 RVM: 0.02
129	9	D: $\ln y = 9.24 + 0.96\ln Q - 0.02\ln Q^2 - 0.13\sin(2\pi\delta t) + 0.09\cos(2\pi\delta t) + 0\delta t + 0.01\delta t^2$ RVM: $\ln y = 9.23 + 0.96\ln Q - 0.02\ln Q^2 - 0.12\sin(2\pi\delta t) + 0.09\cos(2\pi\delta t) + 0\delta t + 0.01\delta t^2$	D: 0.03 RVM: 0.03
131	9	D: $\ln y = 9.21 + 0.86\ln Q - 0.04\ln Q^2 + 0.06\sin(2\pi\delta t) + 0.07\cos(2\pi\delta t) + 0\delta t + 0\delta t^2$ RVM: $\ln y = 9.19 + 0.86\ln Q - 0.04\ln Q^2 + 0.05\sin(2\pi\delta t) + 0.06\cos(2\pi\delta t) + 0\delta t + 0.01\delta t^2$	D: 0.03 RVM: 0.03
150	9	D: $\ln y = 7.58 + 0.86\ln Q - 0.01\ln Q^2 - 0.11\sin(2\pi\delta t) - 0.09\cos(2\pi\delta t) - 0.02\delta t - 0.01\delta t^2$ RVM: $\ln y = 7.55 + 0.87\ln Q - 0.01\ln Q^2 - 0.09\sin(2\pi\delta t) - 0.07\cos(2\pi\delta t) - 0.02\delta t - 0.01\delta t^2$	D: 0.04 RVM: 0.04

159      6      D:  $\ln y = 9.05 + 0.90 \ln Q - 0.04 \ln Q^2 - 0.04 \sin(2\pi\delta t) -$   
 $0.04 \cos(2\pi\delta t)$   
RVM:  $\ln y = 9.04 + 0.91 \ln Q - 0.04 \ln Q^2 - 0.03 \sin(2\pi\delta t) -$   
 $0.02 \cos(2\pi\delta t)$

D: 0.01  
RVM: 0.01

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Table A4. TDS. Estimated covariance matrix of weights at station 122 using RVM

	$a_0$	$a_1$	$a_2$
$a_0$	0.000705	0.000071	-0.000676
$a_1$	0.000071	0.000421	0.000128
$a_2$	-0.000676	0.000128	0.001286

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